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Temporal Variations of Accumulated Cesium in Natural Soils after an Uncharacteristic External Exposure

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Abstract

The accumulation behaviors and solid phase partitioning patterns of stable cesium, which have been recognized as an indicator of the long-term movement of radioactive cesium (^{137}Cs or ^{134}Cs) in ecosystems, were studied in typical and natural soils of Japanese origin, namely, red clay, leaf-mold and andosol soils. The retention and migration of soil-phase cesium have been explained relative to various factors, such as soil organic matter contents, competitive cation concentrations and the adsorption ratio of Cs to the solid phase. Cesium was adsorbed nearly quantitatively in the leaf-mold type soil, and the rate of Cs absorption increased as the particle size decreased in the red clay and andosol soils. The distributions of Cs within the soil solid phases were defined using the selective sequential extraction scheme and were used to explain its relative incorporation in the soil fractions. Solid phase fractionation indicated that nearly half of the total cesium concentrations in the soils were in the ‘residual’ fraction (representing the metal that was incorporated within the crystalline lattice of the soil and was difficult to extract). These findings are expected to provide information regarding suitable conditions for remediation, immobilization or the recovery of cesium from contaminated soils with excess cesium concentrations.

Keywords

Cesium; Accidental exposure; Soil contamination; Solid-phase distribution; Sequential fractionation; Temporal variation

1.0 Introduction

Waste disposal operations or accidental releases due to nuclear-technology-related activities have resulted in the release of large amounts of radionuclides into the environment. The accidental release of radionuclides has been a topic of interest for years, beginning with the catastrophic nuclear accident that occurred on 26 April 1986 at the Chernobyl Nuclear Power Plant [1, 2]. The Fukushima Dai-ichi Nuclear Power Plant (FDNPP) accident that resulted from an earthquake and subsequent tsunami on March 11, 2011 is the most recent example of such a catastrophe. Following the FDNPP accident, a large number (~500) of studies were published that covered the various aspects of the accident. Sorting these published articles by topic in the ISI-listed journals indicated that the studies mainly focused on radioactivity monitoring in places with foods. This trend becomes obvious when considering the importance of estimating the effects of fallout on human health. Approximately 80% of the FDNPP-generated fallout in the contamination zone affected agriculture, including forests [3]. However, very few publications have discussed the effects of fallout on agriculture, including soil contamination due to the extended radionuclide emissions [4-7].

Among the radioactive materials, the dispersion of radiocesium at elevated concentrations evokes concern due to its extended solubility characteristics as an alkaline metal ion, its comparatively longer half-life, and its easy incorporation into living beings [8-10]. The most notable radiocesium isotope is ^{137}Cs , which has a half-life of 30.2 years [11], is abundant in nuclear wastes and in radioactive fallout and can contaminate large areas of agricultural land [12-15]. The FDNPP accident resulted in an estimated atmospheric release of 1.3×10^{16} Bq of ^{137}Cs over four weeks (March 12 and April 6, 2011) [16]. Approximately 10–20% of the emitted radiocesium was deposited over the terrestrial soils in northeastern Japan [15, 17]. The accumulated radiocesium in soils is a threat because plants can take up Cs [18]. Thus, it is important to understand the accumulation and dissemination patterns of radiocesium in the soil to assess the potential health risks of Cs contamination or its impacts on food production after over-exposure occurs. Moreover, radiocesium (particularly ^{137}Cs) is

extremely useful for quantifying soil erosion and deposition risks in forested and agricultural areas where knowledge regarding diffusion processes is important [19, 20].

The fate of radiocesium in the environment and its subsequent migration in soils depends on its physico-chemical form and commonly follows the behavior of its stable form [9, 21, 22]. For example, previous research indicated that the soil-to-plant transfer factor of ^{137}Cs follows a pattern that is similar to stable Cs [22-26]. In addition, analogous distributions of ^{137}Cs have been observed in different rice components relative to stable ^{133}Cs species [27, 28].

The objective of this study is to investigate the cesium distribution in the operationally defined physico-chemical and particle size fractions of soils to understand the temporal variations of cesium after being released at an uncharacteristic rate.

2.0 Experimental

2.1 Instruments

The atomic absorption spectroscopy (AAS) technique was used to determine the stable cesium concentration in solution. An AAnalyst 600 (PerkinElmer, Waltham, MA) was used that was equipped with a transverse heated graphite atomizer with an integrated, pyrolytic graphite coated platform and a longitudinal Zeeman-effect background corrector. The light source was an electrodeless discharge lamp (EDL) that was powered by an EDL System II that was operated at 18 mA. The wavelength was set at the 852.1 nm resonance line and the monochromator spectral bandpass was set at 0.7 nm. In addition, a baseline offset correction time at 2.0 s was used with a read delay of 0.0 s. Argon was used as the purge gas at a flow of 250 mL min⁻¹. The temperature sequence was 110 (drying), 130 (drying), 750 (ashing) and 1900 °C (atomization), which were performed for 30, 30, 20 and 4 s, respectively. The loading volume was 20 µL, with 20 µL of the sample and matrix-modifier (1% H₂SO₄). Calibration was conducted thereby using three standard solutions that contained 0.01, 0.05 and 0.1 mg L⁻¹ cesium.

Inductively coupled plasma optical emission spectrometry (ICP-OES) was used to detect the co-existing trace metals (e.g., Al, Ca, Fe, K, Mg, Mn and Na) in solution. For this

analysis, an iCAP 6300 instrument from Thermo Fisher Scientific (Waltham, MA) was used that included an EMT duo quartz torch, a glass spray chamber and a concentric glass nebulizer. A radio frequency power of 1.15 kW was used at the torch, the plasma, auxiliary and nebulizer gas flows were maintained at 12, 1 and 0.5 L min⁻¹, respectively, and an integration time of 30 s was used.

The microwave-assisted reaction was used to digest the soil samples. For this analysis, a Multiwave 3000 instrument from Anton Paar GmbH (Graz, Austria) was used that was equipped with an 8-position rotor and hydraulic pressurized sensing system for all vessels. The vessels (XF100) supported a controlled pressure of 6 MPa, a maximum operating pressure of 12 MPa and a maximum temperature of 240 °C.

A KDF S-8 muffle furnace from Kenis Scientific (Osaka, Japan), a Digiprep Jr block heater from SCP Science (Quebec, Canada) and a DX 600 oven from Yamato Scientific (Tokyo, Japan) were used for heating or drying.

Ultrapure water with a resistivity > 18.2 MΩ·cm was prepared using the Arium Pro UV water purification system from Sartorius Stedim Biotech GmbH (Göttingen, Germany). The pH measurements were performed using a Navi F-52 pH meter from Horiba Instruments (Kyoto, Japan).

Each instrumental measurement or treatment procedure was performed in three replicates and the resulting averaged value is reported.

2.2 Materials

2.2.1 Soil samples

The soil samples represented three soil types that are common in most regions of Japan (including the Fukushima), red clay, leaf-mold and andosol soils. A large amount (approximately 5 kg) of each soil type was collected from representative locations in Kanazawa, which is located on the Sea of Japan and is bordered by the Japanese Alps. The soil samples were dried for 24 h at 60 °C before separating into size fractions of 2000–212 μm (coarse sand; SF-1), 212–63 μm (fine sand; SF-2), and < 63 μm (silt and clay; SF-3) with the use of grading sieves based on the ISO 14688-1 classification of soils [29].

2.2.2 *Chemicals and laboratory wares*

Analytical grade commercial products were used throughout the study. Cesium nitrate salt and the standard cesium solution (as CsCl) were obtained from Kanto Chemical (Tokyo, Japan) and were used to prepare the working standards. In addition, the HF, HNO₃, H₂SO₄, HCl and NaOH were obtained from Kanto Chemical (Tokyo, Japan) and were used for the decomposition experiments or as mentioned below. Magnesium chloride from Merck KGaA (Darmstadt, Germany), acetic acid from Tama Chemicals (Kanagawa, Japan), hydroxylamine chloride from Wako Pure Chemical (Osaka, Japan), and hydrogen peroxide from Kanto Chemical (Tokyo, Japan) were used during the extraction experiments. The HCl and NaOH solutions (1 mol L⁻¹) were used to adjust the solution pH. Working solutions were prepared from the stock solutions by diluting with ultrapure water on a weight basis.

Low-density polyethylene (PE) bottles from Nalge Nunc (Rochester, NY), screw-capped PE tubes from AS ONE (Osaka, Japan), conical polypropylene centrifuge tubes (Biologix Research, Lenexa, KS), and micropipette tips from Nichiryo (Tokyo, Japan) were used throughout the study. The laboratory equipment was washed by soaking in Scat 20X-PF alkaline detergent from Nacalai Tesque (Kyoto, Japan) overnight and then overnight in 4 mol L⁻¹ HCl before rinsing with ultrapure water after each step.

2.3 *Methods*

2.3.1 *Preparation of Cs-contaminated soil samples*

The granulometric fractions of the dried soils (3 g) were spiked with 30 mL of a standard solution containing 5 mg L⁻¹ cesium (as CsCl). The soil-solution mixtures were continuously shaken in an end-over-end shaker at 200 rpm at 25 °C for 24 h. During this step, the soil particles were completely and evenly exposed to the added cesium. Next, the suspension was centrifuged at 13000 rpm for 30 min to separate the soil solids from the solution using a CR20GIII high-speed refrigerated centrifuge from Hitachi (Tokyo, Japan). The supernatant liquid was carefully separated and analyzed for Cs to calculate the amount of adsorbed Cs in the soil. The residual soil solid fraction was rinsed with the ultrapure water (5 mL × 5) and then dried in an oven at 60 °C for 3 h. The dried Cs-spiked samples were stored

in capped bottles and were stored in the dark at 25 °C until further experiments were conducted.

2.3.2 Determination of the total organic matter content in the raw soils

For all of the *soil* types, 1 g from each of the segregated fractions was dried at 110 ± 5 °C in an oven for 1 h before subjecting the samples to a calcination treatment at 650 ± 50 °C for 2 h in a muffle furnace with an air atmosphere. Next, the ash was cooled and measured to determine the weight loss from the initial value, which indicated the total organic matter content.

2.3.3 Determination of the cation contents in the raw soil

The soil size fractions (0.25 g) were placed in pre-cleaned PTFE vessels in a microwave reaction system with a mixture of HNO₃ (16 mol L⁻¹, 2.5 mL), HCl (12 mol L⁻¹, 1 mL) and HF (28 mol L⁻¹, 1.5 mL). The following operating conditions were used: a maximum unpulsed microwave power of (P_{\max}) = 1400 W; a ramp time of (R_T) = 5 min; a hold time of (H_T) = 35 min at maximum power; and a vent time of (V_T) = 20 min. The soil-liquid suspensions from the digestion procedure were filtered using cellulose membrane filters with a pore size of 0.45 μm from Advantec (Tokyo, Japan). Next, the filtrate was evaporated on a block heater assembly at 60 °C until the total volume was reduced to a few milliliters. Next, the concentrated solution were diluted to a volume of 50 mL using ultrapure water and were analyzed for their total Al, Ca, Fe, K, Mg, Mn and Na contents using ICP-OES.

2.3.4 Mapping of the solid-phase cesium distributions in the soil

The selective sequential extraction (SSE) scheme, which was originally proposed by Tessier, *et al.* [30], was used to map the mobility and distribution of cesium in the soil-solid phase. The scheme was based on washing the target soil sample with a series of gradually aggressive extractants [30-32]. The soil cesium was divided into five different fractions, exchangeable, acid soluble or bound to carbonates, reducible or bound to iron and manganese oxides, oxidisable or bound to the organic matter, and residual. The compositions of the extractants that were used to delineate the distinctive operationally defined fractions

are shown in Table 1. Hereafter, the fractions are referred to as F1, F2, F3, F4 and F5, respectively.

2.3.5 Determination of the soil cesium contents

The pre-cleaned PTFE vessels of the microwave reaction assembly were fed with 0.25 g of the raw, Cs-spiked, and SSE scheme residue soils before the sequential addition of 14 mol L⁻¹ HNO₃ (1 mL), 12 mol L⁻¹ HCl (1 mL) and 28 mol L⁻¹ HF (1 mL). The acid decomposition was concluded using a three-step microwave-assisted reaction. The first and second steps were conducted consecutively under the following conditions: (a) $P_{\max} = 500$ W, $R_T = 10$ min, and $H_T = 15$ min; and (b) $P_{\max} = 1400$ W, $R_T = 10$ min, $H_T = 60$ min, and $V_T = 15$ min. In the third step, 5% H₃BO₃ (10 mL) was added to the soil-acid mixture under the following conditions: $P_{\max} = 1400$ W; $R_T = 0$ min; $H_T = 15$ min; and $V_T = 15$ min. The three-step processes ensured the complete dissolution of the soil solids. The remaining solution from the PTFE vessel was carefully collected, diluted to 50 mL using ultrapure water, and analyzed for total Cs using GF-AAS.

3.0 Results and discussion

3.1 Effects of soil type and soil particle size on the Cs assimilation in the soils

The natural concentrations (mg kg⁻¹) of stable cesium (¹³³Cs) in the leaf-mold, andosol and red clay soil types were 27.9 ± 3.2 , 23.0 ± 1.4 , and 6.5 ± 1.3 , respectively (Figure 1). These values corresponded to the estimated cesium concentrations in natural soils of between 0 to 26 mg kg⁻¹ [8]. The natural cesium concentrations in the red clay soil occurred in the finer particle size fraction (i.e., silt and clay; SF-3). In contrast, for the leaf-mold and andosol soil types, the distribution was uniform with 7 to 11 mg kg⁻¹ of cesium in all of the soil size fractions, including the coarse sand (SF-1) and fine sand (SF-2) fractions. The ¹³³Cs distribution patterns and the size fractions of the red clay soils agree with previous findings regarding radiocesium, which indicated that the ¹³⁷Cs contents in the soil samples were associated with the finer size fractions and exhibited greater activities in the finer size fractions than in the bulk samples [9, 33, 34]. The adsorbed cesium concentrations (mg kg⁻¹) were 103 ± 2.5 , 116 ± 3.9 and 121 ± 5.4 in the leaf-mold, andosol and red clay soils,

respectively, when atypically exposure occurred from an external source. In addition, the adsorption data indicated the possibility of higher cesium retention rates in the red clay soil under sudden extreme exposures relative to the other soil types. However, irrespective of the soil type, the total cesium content within the soil particle size fractions was more or less even and was in the range of 40 to 48 mg kg⁻¹ (Figure 1). This pseudo-uniformity in the cesium distribution in the soil particle size fractions can be attributed to the shorter time-lag between the exposure period and concentration measurements [35]. The particle size can contribute to several effects when ¹³⁷Cs is used to estimate soil erosion rates. Furthermore, the particle size effects can be accounted for by using conversion models. However, these effects were not evident in other studies. If no apparent enrichment is detected in soils that were exposed to high activities, a particle size correction may not be needed for the conversion models [20, 36-39].

3.2 Effects of soil organic matter contents on Cs assimilation in soils

The adsorption of cesium in soils decreases as the amount of soil organic matter increases, which implies that cesium is more available to biological systems in soils with high organic matter contents [40, 41]. This hypothesis is somewhat supported from the lowest adsorption ratio of cesium in the leaf-mold soil. The leaf-mold soil type is defined by humus or compost that consists of decomposed leaves and other organic materials. A gradual increase in the soil organic matter content was observed among the coarse sand (SF-1), fine sand (SF-2), and silt and clay (SF-3) size fractions (Figure 2a). However, the relationship between the soil cesium content (mg kg⁻¹) and the organic matter (%) is rather discrete (Figure 2b).

3.3 Effects of coexisting cation concentrations on Cs assimilation in soils

The interactions of metal cations with mineral surfaces in and near hazardous chemicals and radioactive waste sites often control the mobility of alkali metals and their radioisotopes [42, 43]. The Al, Ca, Fe, K, Mg, Mn and Na concentrations in raw soils have been measured in different soil size fractions (Figure 3). In each soil type, a greater K⁺ concentration was observed in the coarse sand fraction, while the Ca²⁺ and Mg²⁺ contents were higher in the fine

sand fractions. The Na^+ ion contents were comparatively lower in the different Japanese soil types, and the distribution patterns in the different size fractions were similar to those of K^+ .

The sorption of cations in soil depends on ion-exchange processes, while the selectivity and fixation of the cations depend on the respective hydration energies. Cations with low hydration energies, such as K^+ , are fixed in the interlayers of the soil particles and are attributed to interlayer dehydration and layer collapse. In contrast, cations with high hydration energies (e.g., Ca^{2+} , and Mg^{2+}) create expanded interlayers and are not fixed [44-46]. Furthermore, the Cs^+ cation has a low hydration energy, and the selection of Cs^+ over Ca^{2+} and Na^+ ions have been reported in soils [45-47]. The differences in the Ca^{2+} , Mg^{2+} and Na^+ ion fixation patterns relative to K^+ within the soil interlayers indicate that K^+ may negatively influence Cs^+ retention patterns in the soils.

The soil Al^{3+} concentrations were larger than the concentrations of other ions. In addition, the Fe^{3+} concentration was moderate, and the Mn^+ concentrations were insignificant for all of the soil types. The cations with low hydration energies were retained more strongly and were interstratified regularly, likely due to the oxidation of Fe ions within the soil layers [48]. In addition, the Al^{3+} or Fe^{3+} coatings on the soil particles resulted in higher cation sorption [49]. Thus, stronger Cs^+ retention was expected on the red clay, leaf-mold and andosol soils. Furthermore, the combined Al^{3+} and Fe^{3+} concentrations were much higher than the K^+ concentrations in each soil type and soil size fraction, which likely suppressed any retention or hindering impacts.

3.4 The distribution of solid-phase Cs in soils: Similarities between ^{133}Cs and ^{137}Cs

The SSE scheme uses several extracting solutions in sequence, and the varying reactivities and behaviors of the soil-bound metals toward the extracting solutions are correlated with the metal partitioning within the solid phases of the soils [50]. The operationally defined fractionation of ^{133}Cs in the Cs-loaded Japanese soils and the impacts of the Cs distributions over 6 months are shown in Figure 4.

The SSE treatment revealed that 90, 83.7 and 84.9% of the Cs contents in the red clay, leaf-mold and andosol soils, respectively, were incorporated in the non-detrital fractions during the earlier stages of exposure, while residual fraction contents varied from 10 to 16%

(Figure 4a). The cumulative Cs contents in the exchangeable (F1), acid-soluble (F2), reducible (F3) and oxidisable (F4) fractions represented the ‘non-detrital’ fraction with mobile characteristics. The Cs fraction that was incorporated within the crystalline lattice of the soil with persistent environmental interactions was associated with the residual fraction (F5) [32, 51-53]. The temporal Cs-distribution data indicated a common trend of decreasing Cs concentrations in the F1 fraction. However, the incorporation gradually increased in the F5 fraction. The negative variations of the F5 Cs content after one month of exposure were deliberately ignored because the period was too short to result in any convincing movement patterns within the soil phases. After 5 months, the Cs concentration in the F2 fraction increased, while the Cs concentrations in the F3 and F4 fractions decreased (Figure 4a). A similar temporal trend was observed for the different soil size fractions, regardless of the soil type (Figure 4b). The Cs concentrations in the different size fractions decreased in descending order in the leaf-mold and andosol soils and were greater in the B size fraction of the red clay soils. The data patterns from the soil size fractions indicated that the increases in the F2 and F5 Cs contents with time were not correlated with the decreasing Cs contents of the particular solid phase Cs contents (F1, F3 or F4). In addition, a mutual internal suppleness could be a more acceptable assumption for interpreting this behavior.

The metal loading in the soils due to anthropogenic affects were limited to the ionic forms or sorbed metals. However, the metals may also exist in their particulate or low solubility forms [54-57], and a random distribution of particles that are rich in metals may appear in the soil rather than the sorbed state. Sequential extraction methods have often been used to evaluate the bioavailability and mobility of radionuclides in soils based on the distribution of the physico-chemical forms of radiocesium. The solid-phase fractionation patterns of ^{133}Cs -loaded and simulated Japanese soils follow fractionation patters that are similar to ^{137}Cs , with a higher rate of accumulation in the residual soil fraction [9, 58-60]. The ^{133}Cs content can reach up to 94% of the total ^{137}Cs in the soils that were exposed to radiocesium contamination from the FDNPP accident [61]. The effects of elapsed time following the contamination of soils with ^{137}Cs were observed during our study of the ^{133}Cs contaminated samples [9, 59, 62]. The lower hydration energies of the ionic cesium resulted

in the augmented occurrence of cesium in the interlayer or frayed edge sites of the soil particles [45-47]. In addition, the incorporation of radiocesium or stable cesium species in the most persistent fraction in the soil solid phases increased.

The kinetics of cesium sorption and the apparent reversibility of the sorption processes are intra-related. These processes are influenced by the slow sorption kinetics, competing cation characteristics, or both [63, 64]. Thus, it was assumed that cesium migrates slowly to energetically favorable interlayer sites. Once the cesium reaches these sites, it is not easily released [64, 65]. Based on the above discussion and the “three-box” model that was proposed by Comans and Hockley [65], a schematic diagram (as illustrated in Figure 5) was used to summarize the temporal variations of the Cs accumulation patterns in the soils after uncharacteristic exposure.

4.0 Conclusions

The distribution of Cs in soils that were exposed to ^{133}Cs (red clay, leaf-mold and andosol soils) was determined in different operationally defined physico-chemical and particle size fractions. A pseudo-uniformity in the Cs distribution was observed in the soil particle size fractions regardless of soil type. In addition, the red clay soil was more vulnerable to Cs during the uncharacteristic high-dose exposure. Greater soil organic matter contents resulted in lower Cs retention rates. Furthermore, the soil cation concentrations affected the rate of soil Cs uptake. The incorporation of Cs mainly occurred in the strongly bound fraction of the soil solid phase, and the temporal movement enhanced the content in the same fraction. Therefore, the long term Cs migration pattern indicated that the movement of the initially mobile Cs fraction relative to the irreversible soil core resulted in a Cs fraction that was less labile for soil-to-plant transfer.

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Table 1: Chemical interpretation and the extraction conditions of the protocol that was used for fractionating the solid phase cesium from the soil

| Fraction | Chemical interpretation | Extraction conditions (for 0.25 g soil) | | | |
|----------|---------------------------|--------------------------------------------------------------------------------------------------------------------|---------------|-------------|--------------|
| | | Extractants and other conditions | Duration h | Temp. °C | Agitation |
| F1 | Exchangeable | 1 M MgCl ₂ (pH 7; 4 mL) | 1 | 25 ± 2 | continuous |
| F2 | Acid soluble ^a | 1 M NaOAc (pH 5, adjusted with acetic acid; 4 mL) | 5 | 25 ± 2 | continuous |
| F3 | Reducible ^a | 0.04 M NH ₂ OH·HCl in 25% (v/v) acetic acid (10 mL) | 6 | 96 ± 3 | occasional |
| F4 | Oxidisable ^a | 0.02 M HNO ₃ (3 mL) + 30% H ₂ O ₂ (pH 2, adjusted with HNO ₃ ; 2.5 mL) | 2 | 85 ± 3 | occasional |
| | | 30% H ₂ O ₂ (pH 2, adjusted with HNO ₃ ; 1.5 mL) | 3 | 85 ± 3 | intermittent |
| | | 3.2 M NH ₄ OAc in 20% (v/v) HNO ₃ (2.5 mL) | 0.5 | 85 ± 3 | continuous |
| F5 | Residual | Microwave-assisted acid decomposition ^b | – | – | – |

^a The fraction names ‘acid-soluble’, ‘reducible’ and ‘oxidisable’ were originally called ‘bound to carbonates’, ‘bound to iron and manganese oxides’ and ‘bound to the organic matter’ by Tessier, *et al.* [30].

^b The detailed microwave-assisted acid decomposition procedure for determining the soil cesium contents is available in Section 2.3.5.

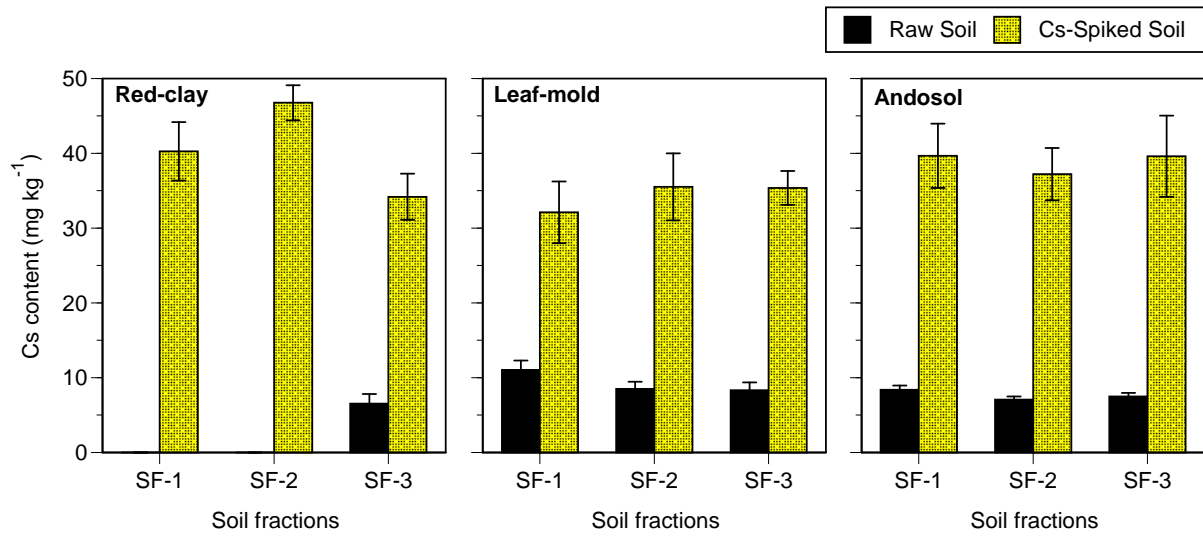


Figure 1: Cesium concentrations in the raw and cesium-spiked soils ($n = 3$). SF-1: Size fraction of 2000–212 μm (coarse sand); SF-2: Size fraction of 212–63 μm (fine sand); and SF-3: Size fraction of < 63 μm (silt and clay).

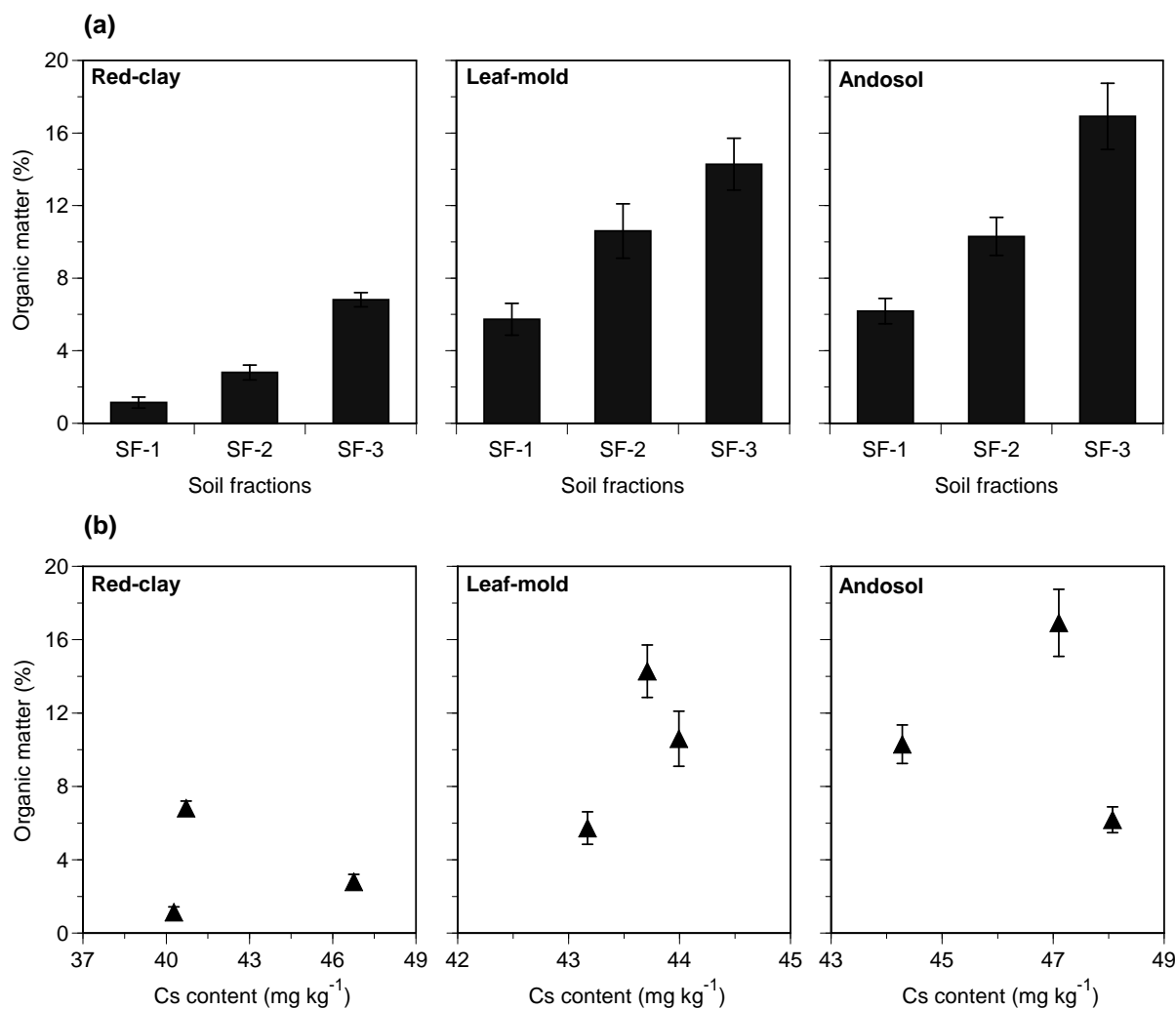


Figure 2: (a) The distribution patterns of the organic matter contents in the soils. (b) The correlation between the cesium accumulation pattern and the organic matter contents in the soils ($n = 3$). SF-1: Size fraction of 2000–212 μm (coarse sand); SF-2: Size fraction of 212–63 μm (fine sand); and SF-3: Size fractions of $< 63\mu\text{m}$ (silt and clay).

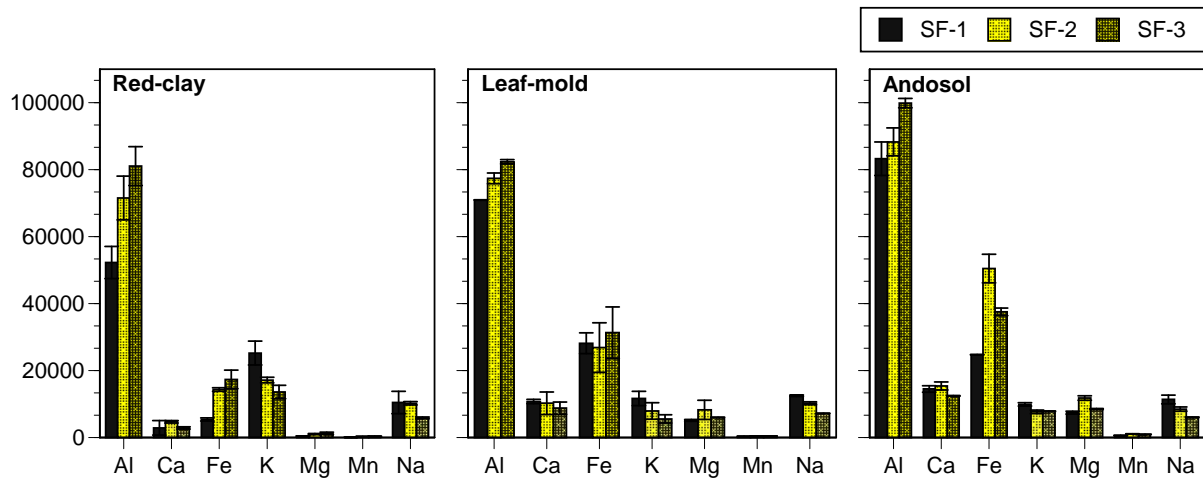


Figure 3: The distribution pattern of the coexisting soil cations ($n = 3$). SF-1: Size fraction of 2000–212 μm (coarse sand); SF-2: Size fraction of 212–63 μm (fine sand); SF-3: and Size fraction of < 63 μm (silt and clay).

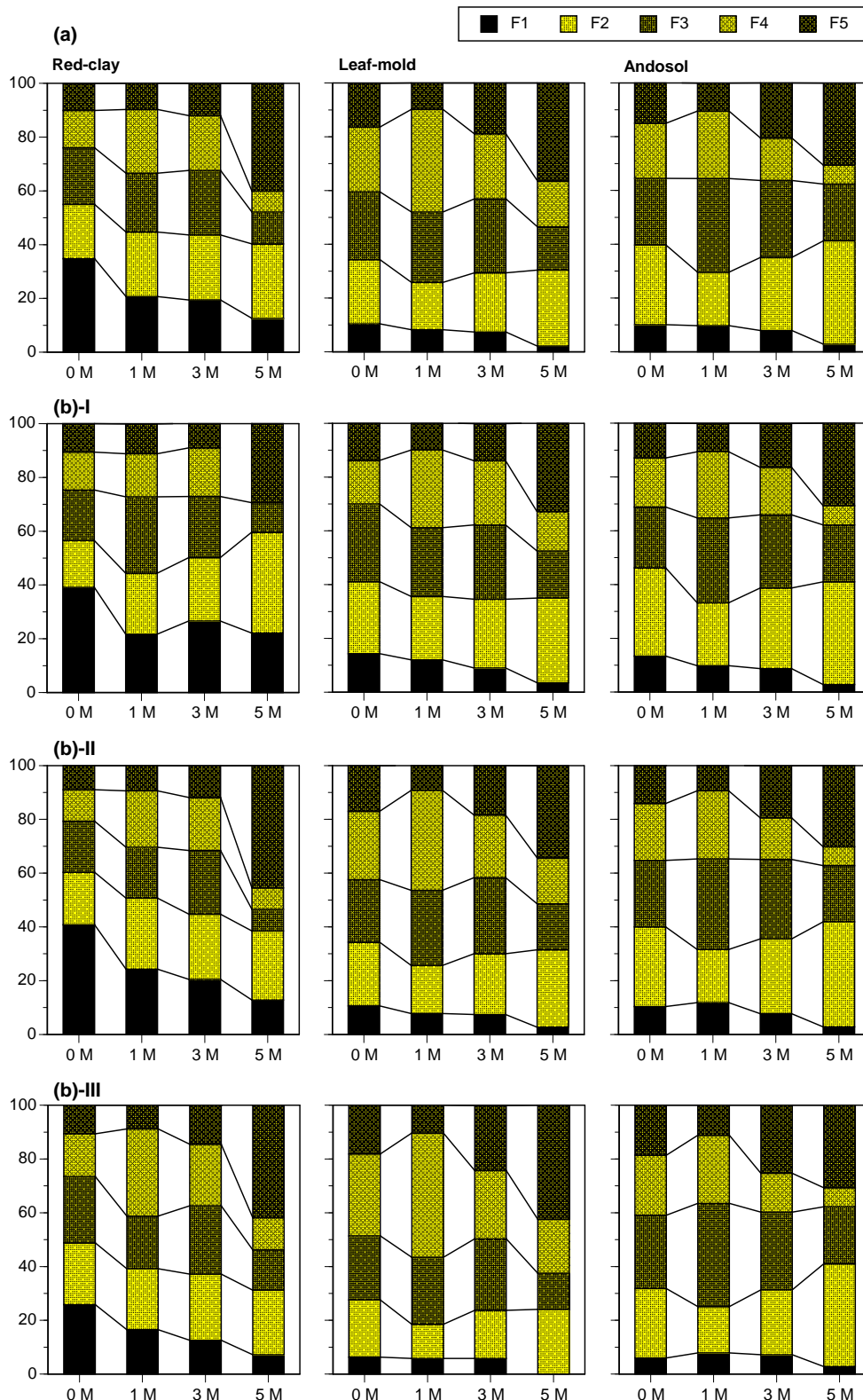


Figure 4: Comparative temporal distribution of cesium in the soil solid-phases: (a) without size fractionation, (b-I) size fraction of 2000–212 μm (coarse sand), (b-II) size fraction of 212–63 μm (fine sand), and (b-III) size fraction of < 63 μm (silt and clay). F1: Exchangeable, F2: Acid soluble, F3: Reducible, F4: Oxidisable, and F5: Residual.

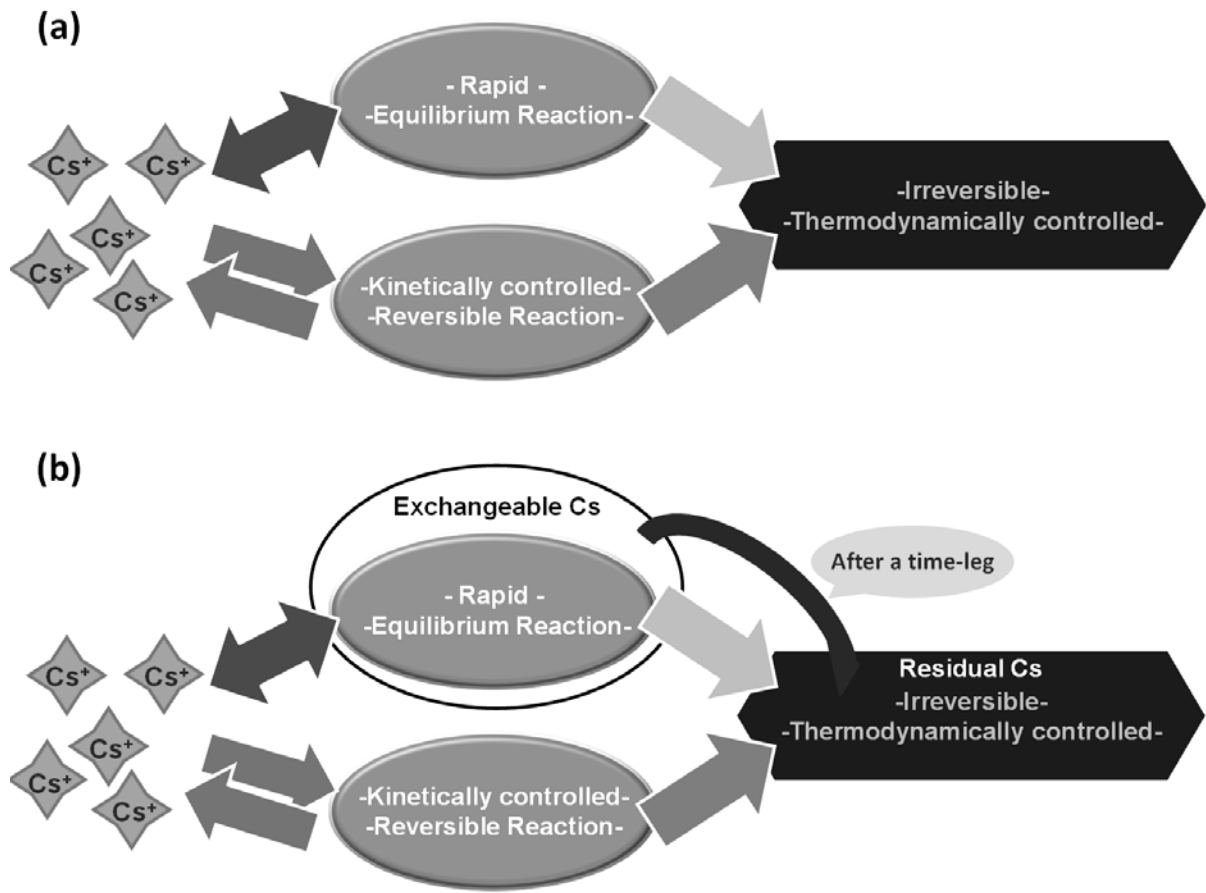


Figure 5: A schematic model of (a) the instantaneous Cs-sorption patterns in the soils and (b) the temporal changes of the Cs-sorption patterns in the soils.